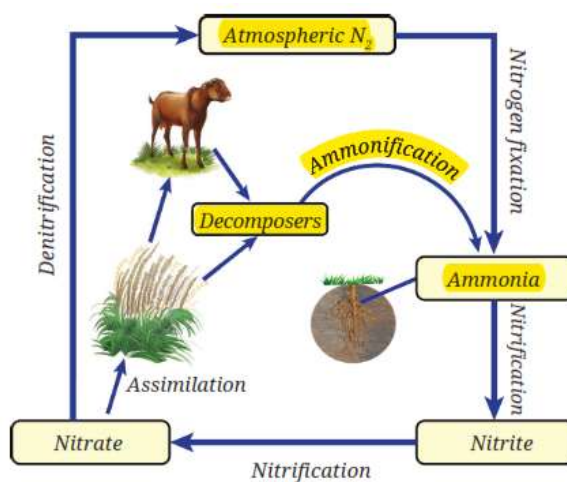
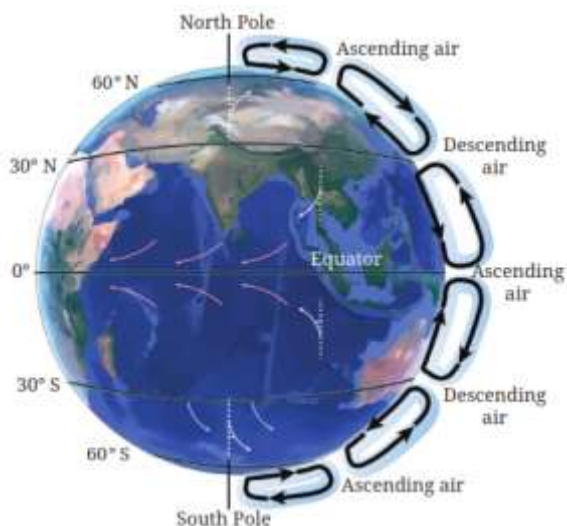
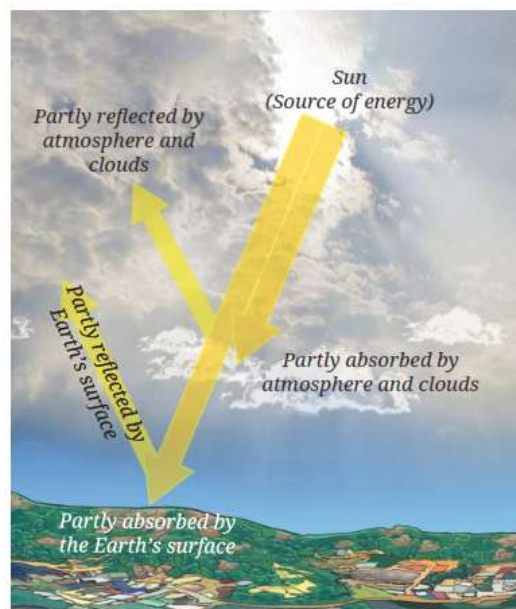


Earth as a System

Energy, Matter, and Life

Chapter Outline

- ✧ Uneven Heating on the Earth
- ✧ Winds and Ocean Currents
- ✧ Biogeochemical Cycle
- ✧ Human Impact on Earth's Processes



INTRODUCTION

In this chapter, we will consider these processes together as belonging to one Earth system made up of interacting 'spheres' -

- Geosphere: Solid rocks, soil, landforms (like the Deccan plateau and the Thar desert), and the Earth's interior.
- Hydrosphere: Liquid water in the form of surface water, such as oceans, rivers (like the Ganga-Brahmaputra river system), lakes and groundwater.
- Cryosphere: Solid form of water, such as ice and snow (like the Himalayan glaciers, snow in Ladakh and polar ice caps).
- Atmosphere: The air surrounding the Earth that we breathe (cleaner air in the mountains and forests).
- Biosphere: All living organisms and their habitats (including mangroves, forests, farms, ocean plankton and coral reefs).

We will explore how energy and matter move, interact across the Earth's spheres, and how a change in one affects others. Understanding these processes has been an important part of the continuing journey of the exploration of science. Natural processes, such as heating by solar radiation, movement of air and water, and nutrient cycling connect these spheres in a delicate balance.

Uneven Heating of the Earth

Solar radiation is the main source of energy on the Earth. It reaches the Earth as electromagnetic (EM) waves that travel through a vacuum at the speed of light (unlike EM waves, sound waves, which you studied in Chapter 10, Sound Waves: Characteristics and Applications, are mechanical waves and require a medium to travel). The speed of light in vacuum is $3 \times 10^8 \text{ ms}^{-1}$. EM waves cover a wide range of frequencies, from high frequency or short wavelength radiation (gamma rays and X-rays) to low frequency or long wavelength radiation (infrared and radio waves). The high frequency EM waves, such as gamma rays and X-rays have very high energy and can be harmful for life on Earth.

The entire range of electromagnetic radiation called electromagnetic spectrum is shown in Fig. However, the solar radiation reaching Earth is concentrated mainly in the ultraviolet (UV), visible and infrared (IR) range-about 99 per cent of the Sun's energy falls within these wavelengths. These three regions of the spectrum shape the Earth's climate and support life. Gamma rays and X-rays are mostly filtered by the Earth's upper atmosphere, while microwaves and radio waves carry very little energy to significantly warm the Earth.

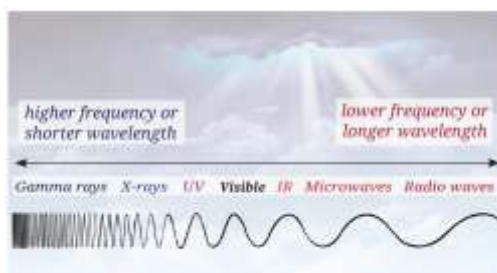


Fig.: Electromagnetic spectrum

Short-wavelength UV radiation is mostly absorbed by the ozone layer in the upper atmosphere, protecting life and contributing to some atmospheric heating. Visible light from the Sun reaches the Earth's surface and provides energy for photosynthesis, which is the primary source of food for most organisms. It also partly warms the land and water. Infrared radiation warms the Earth's surface, which then re-radiates this heat back into the atmosphere. The greenhouse gases, such as carbon dioxide, trap a portion of this outgoing heat (CO_2), methane (CH_4) and water vapour, keeping the Earth warm enough to support life.

Note:

The UV rays lie in the wavelength range of 100 nm to 400 nm [1 nanometre (nm) = 10^{-9} m] and have a much higher energy than the visible light. Prolonged exposure to these rays can damage the eyes and skin, and increase the risk of cancer. Therefore, people should use UV protective glasses and sunscreen if they are likely to be exposed to UV rays. UV rays are also useful in killing germs in water purifiers and help power fluorescent lights.

The amount of the Sun's radiation that reaches the Earth's surface is called insolation. It is responsible for warming the Earth's surface and its atmosphere. The average amount of solar energy received per unit time per unit area, perpendicular to the Sun's rays at the top of Earth's atmosphere, is called the solar constant. Its value is approximately 1.4 kilowatts per square metre (1.4kWm^{-2}), or about 1400 joules per second per square metre ($1400\text{ J s}^{-1}\text{ m}^{-2}$). It represents the Sun's energy available on the Earth before any absorption, scattering or reflection occurs in the atmosphere.

The solar constant is important because it helps scientists understand the Earth's energy balance, climate and weather patterns. However, some of this energy is absorbed and scattered by gases, clouds and dust particles in the atmosphere before reaching the surface of the Earth. As a result, the maximum insolation reaching the Earth's surface is lower than the solar constant and is about 1kWm^{-2} under clear sky conditions. India, due to its geographical location in the tropical and sub-tropical regions, receives abundant sunlight throughout the year. This makes solar insolation highly significant as it drives the southwest monsoon, which influences India's climate and agriculture. It also offers immense potential for harnessing solar energy as a renewable and sustainable source of energy.

Science and Society

Anna Mani, India's pioneering atmospheric scientist, mapped solar insolation across India in the 1950s. She later published *Solar Radiation Over India* in 1981, creating the country's first insolation atlas. Her pioneering measurements showed India's vast solar energy potential. Today, this is being realised through the large-scale deployment of solar power across the country. India is laying the foundation for a resilient, sustainable and solar-powered energy future that sets global benchmarks.

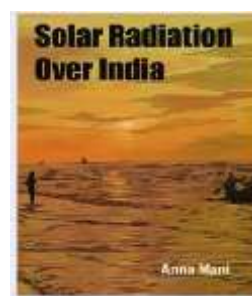


Fig.: Solar Radiation Over India, a book by Anna Mani

Example 1: How much solar energy will be received by a 1 m^2 area in one hour, if the insolation on the surface of the Earth were 1kWm^{-2} ?

Answer: $E = \text{Intensity} \times \text{area} \times \text{time}$

$$E = 1 \times 1000\text{ J s}^{-1}\text{ m}^{-2} \times 1\text{ m}^2 \times 3600\text{ s} \Rightarrow E = 3600000\text{ J} = 3.6 \times 10^6\text{ J}$$

(This amount of energy is approximately equal to the energy needed to melt 5 kg of ice and heat the water obtained to 100°C . It is also equal to one unit of electricity used in a household, as shown on an electricity bill).

Interaction of solar radiation on the Earth's surface

Different materials absorb and heat up differently under sunlight. You have already learnt how land heats up faster than water. There may be some variation due to the material forming the land and the colour of the soil. Dark surfaces absorb more sunlight, while light-coloured surfaces reflect more and remain comparatively cooler. For example, dark coloured roads heat up more quickly, while light-coloured surfaces remain comparatively cooler. You have also learnt why dark-coloured clothes feel hotter than white ones in summer.

The fraction of solar radiation reflected by a surface is called its albedo (the word 'albedo' comes from Latin, which means whiteness). High albedo surfaces stay cool because they reflect more light, while low albedo ones heat up more quickly as they reflect less and absorb more light



Fig. Solar radiation and warming of the Earth

Table: Reflection of solar radiation by surfaces of materials

S. No.	Materials	Albedo
1.	Snow	0.80-0.90
2.	Ice	0.50-0.70
3.	Crushed rock	0.25-0.30

You must have seen that snow and ice have high albedo, meaning they reflect a large proportion of the incoming solar radiation. This makes polar regions very cold. On the other hand, surfaces like black soil and ocean water have lower albedo as they absorb more solar radiation, and hence, they are relatively warmer.

We feel hot inside a concrete house at night especially during hot summer, this is due to the heat re-radiated by the concrete during the night. However, traditional houses built with thick mud and wooden walls offer cool conditions even in hot summer due to less re-radiation.

Urban Heat Island Effect :

Cities are warmer than their surrounding rural areas, especially during summer and at night. Compared to rural areas, cities have more built-up areas, including buildings of steel, concrete, brick, and roads made up of concrete and asphalt. All these materials absorb solar radiation and retain heat. The re-radiated heat from these materials warms cities more than the surrounding rural areas. This increases energy demand for air conditioning, further stressing the urban ecosystems. Rural areas and forests have more vegetation and stay cool through shade and plant transpiration. The urban heat island effect shows how human land use can alter the local climate



Fig. Urban heat island effect in a city

The amount of radiation received by a particular region on the Earth's surface also depends on its latitude. The uneven heating due to the latitude of a place is important as it determines the atmospheric conditions of that place.

Latitude and Earth's shape

You have already learnt in earlier grades that the Earth is spherical. As a result, the Sun's rays strike different latitudes at different angles on the Earth. The Sun's radiation falling on the equatorial region is concentrated over a smaller area, whereas it is spread over a larger area in the polar region. This is why equatorial regions remain relatively warm throughout the year, whereas polar regions experience much colder conditions. This uneven heating creates temperature differences between the equator and the poles.

You have also learnt how the Earth's spherical shape and the tilt of the Earth's axis of rotation give rise to seasons, and the changing length of daytime during one revolution of the Earth around the Sun. Thus, solar radiation is not evenly distributed across the globe. This uneven heating of the Earth's surface across the globe drives global winds and ocean currents.

Role of the atmosphere

The air surrounding the Earth is called the atmosphere. It is held in place by the force of Earth's gravity. It consists mainly of nitrogen (78%) and oxygen (21%) along with small amounts of argon, carbon dioxide, water vapour, and other gases.

The Earth's atmosphere has a layered structure. This structure of the atmosphere is important as it helps explain key weather patterns (Fig.). Nearly all weather phenomena take place in the troposphere (with an average height of about 12 km), which is heated from the Earth's surface. In this layer, temperature decreases with height ($\sim 6.5^{\circ}\text{C}/\text{km}$). As the warm air rises, it drives winds and storms. The height of the troposphere is maximum above the equator and lowest above the polar regions.

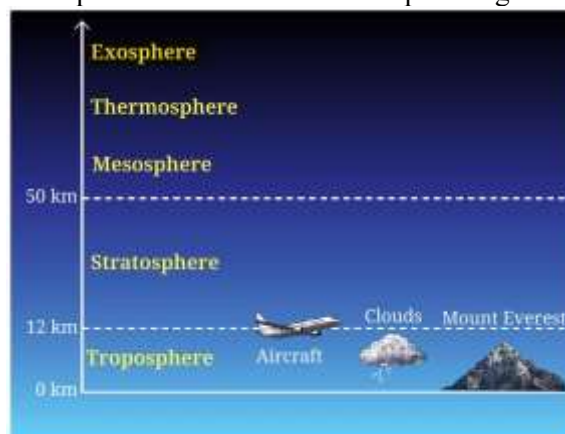


Fig. Layers of the Earth's atmosphere

In the stratosphere (12 – 50 km), where the ozone layer is located, ozone absorbs UV rays and heats up the atmosphere, so temperature rises with height. This increase in temperature with height in the stratosphere calms the layer due to the lack of vertical mixing of air, keeping weather confined to the troposphere. Above the stratosphere are the mesosphere, thermosphere and exosphere, but they play only a minor role in regulating climate on the surface of the Earth. At about 100 km above the Earth is the start of the region we call outer space. These atmospheric layers show how the atmosphere regulates energy flow, shaping weather, climate and life.

Table: Key atmospheric layers

Layer	Approximate altitudes	Features
Troposphere	0 – 12 km	Weather formation; temperature decreases with height
Stratosphere	12 – 50 km	Ozone layer absorbs UV; temperature increases with height

The atmosphere plays two crucial roles in protecting life on Earth. Firstly, it partly absorbs the incoming solar radiation. The ozone layer blocks the harmful UV rays. Clouds and other gases also absorb some sunlight before it reaches the surface of the Earth (Fig.). Secondly, it traps outgoing heat. The Earth's surface absorbs sunlight and re-radiates it in the infrared region. Greenhouse gases like CO_2 , CH_4 and water vapour absorb this re-radiated heat, preventing it from escaping into space. Without the atmosphere, the Earth would be too cold for life to survive. However, excess CO_2 from human activities enhances the greenhouse effect, causing global warming, which if left unchecked, could make the Earth uninhabitable.

Venus is hotter than Mercury, even though Mercury is closer to the Sun. This is because of the presence of an atmosphere on Venus, which has led to an uncontrolled greenhouse effect.

The atmosphere is continuously modified by natural processes, influencing weather, climate, and energy balance. When solar radiation and cosmic rays interact with the atmosphere, they influence weather and climate, and affect communication systems. Thus, the atmosphere makes the Earth a unique planet.

Why is the ozone layer so important?

The ozone layer is vital to life on the Earth because it acts as a protective shield, absorbing harmful UV radiation from the Sun. When ozone molecules are destroyed faster than they are naturally formed, the layer becomes thinner and less effective. In the late 20th century, human-made chemicals called chlorofluorocarbons (CFCs), used in refrigerators and aerosols, caused severe ozone loss over Antarctica, which came to be known as the ozone hole. Increased UV radiation on the Earth can harm the living organisms and the ecosystems. A global agreement known as the Montreal Protocol reduced the use of CFCs and the ozone layer is now slowly recovering, showing the power of international scientific cooperation.

Uneven Heating Causes Wind and Ocean Currents

You have learnt that wind is the movement of air from a region of high pressure to a region of low pressure. These pressure differences are mainly caused by the uneven heating of the Earth's surface by the Sun. Uneven heating of land and water is also responsible for phenomena such as land and sea breezes, which you have studied earlier. The same basic idea helps us understand the formation of winds and ocean currents on vastly different scales.

Local winds

Uneven heating of the Earth's surface also produces local winds, such as valley and mountain breezes. In mountainous regions, the slopes and the valley floor do not heat up and cool down at the same rate. During the day, the mountain slopes facing the Sun are heated more rapidly than the valley floor. As a result, the air over the slopes becomes warm and rises, creating a low-pressure region. Cooler air from the valley moves up the slopes to replace the rising warm air. This flow of air is called a valley breeze (Fig. a).



Fig. (a): Valley breeze

After sunset, the situation reverses. The mountain slopes lose heat faster and become cooler, while the valley floor remains relatively warmer. The air over the slopes becomes cooler and denser, and flows down into the valley. This is known as a mountain breeze (Fig.b). Such daily changes in the direction of the wind are commonly experienced in hilly regions like Shimla, Dehradun, and other Himalayan valleys. These local winds influence the weather conditions, agriculture, and daily life in mountainous regions. They help regulate temperature and moisture conditions and support soil and crop health. Thus, uneven heating of the Earth's surface sets up winds and forms an important part of the Earth's atmospheric circulation.



Fig. (b): Mountain breeze

Planetary winds

Uneven heating of the Earth between the equator and the poles creates belts of low and high pressure. This large-scale pressure difference sets the air in motion over long distances, giving rise to planetary winds.

Near the equator, intense heating by the Sun causes warm air to rise, forming an equatorial low-pressure belt. As this air rises, it moves poleward at higher altitudes. On cooling, it becomes denser and sinks around 30° North and South latitudes, forming the sub-tropical high-pressure belts. From these high-pressure regions, air flows back towards the equator along the Earth's surface, completing one circulation cycle (Fig. a). However, not all the sinking air at the sub-tropical belt returns to the equator. A part of it moves towards the poles along the surface, and rises again around 60° North and South latitudes, where it meets the cold air from the polar regions. This creates the sub-polar low pressure belts.

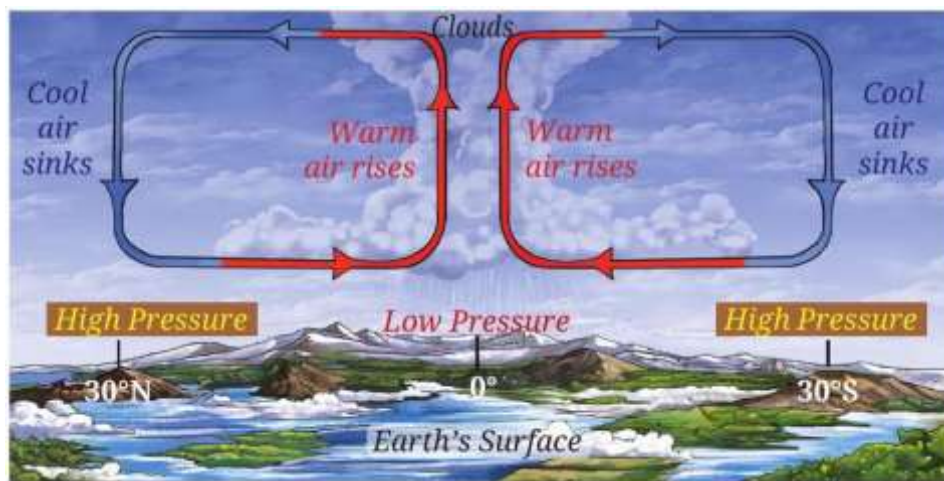


Fig. (a): Wind circulation between equatorial low pressure belt and sub-tropical high pressure belt

In the polar regions (around 90° North and South), very low temperatures cause cold and dense air to sink, forming polar high pressure belts. Air from these regions flows towards the sub-polar belts, completing another circulation cycle (Fig. b). The Earth's rotation causes these winds to be deflected from their straight paths. As a result, planetary winds follow curved paths rather than moving directly from high-pressure to low-pressure regions. The winds are deflected towards the right in the Northern Hemisphere and towards the left in the Southern Hemisphere due to the rotation of the Earth on its axis.

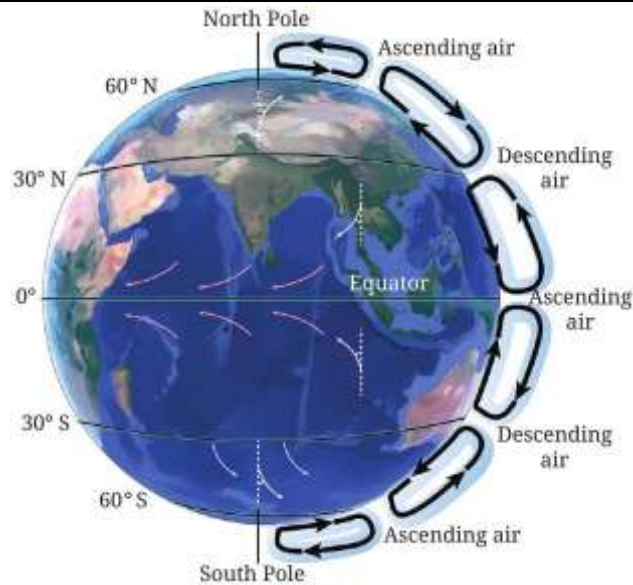


Fig. (b): The distribution of pressure belts and planetary winds

Ocean currents

Ocean currents are the continuous movement of large masses of ocean water. Similar to winds, planetary pressure differences also drive the ocean currents. Strong planetary winds drag the surface water of the oceans because of friction, setting surface currents in motion. In addition to winds, differences in temperature and salinity, the rotation of the Earth, and the distribution of land masses influence the movement of ocean water. The water of the equatorial and polar regions has different temperatures. While warm equatorial waters travel over the surface toward the poles, colder and denser waters slowly flow back towards the equator through deeper ocean levels. Ocean water has different salinities depending on the location. Water with lower salinity, being less dense, tends to remain near the surface, while higher salinity, denser water sinks and moves at deeper levels in the ocean. The Earth's rotation deflects these moving water masses, forming large circular patterns called gyres. These rotate clockwise in the Northern Hemisphere and counterclockwise in the Southern Hemisphere (Fig. 1.10a). Continents further modify these paths by blocking and redirecting currents.

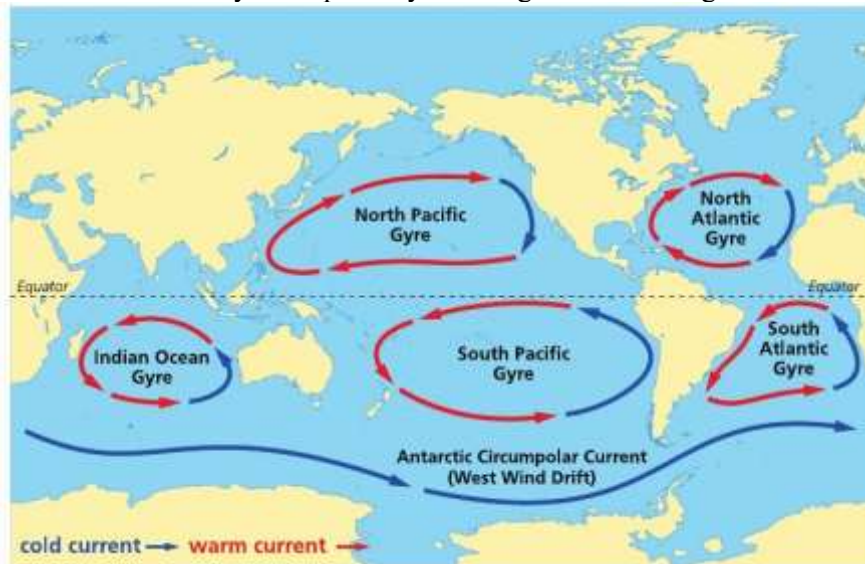


Fig. (a): Global surface ocean currents showing circulation like gyres

Ocean currents play a major role in regulating the Earth's climate and supporting life. By transporting heat from the equator towards the poles, they reduce temperature differences across the planet. For example, the warm ocean current, the North Atlantic Drift (Fig. b), an extension of the Gulf Stream (a strong ocean current carrying warm water from the southern part of the North American east coast across the Atlantic Ocean), flows toward the northwestern coast of Europe and keeps many ports ice-free during winter, even at high latitudes. This moderating influence on climate also supports human activities, such as trade and commerce. Ocean currents also support a massive ecosystem by transporting nutrients.

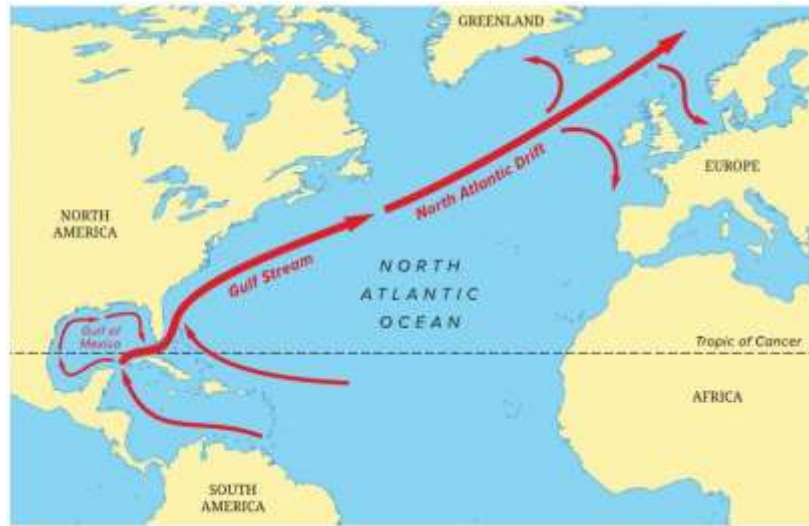


Fig. (b): Gulf Stream brings warm water from the equator

India's Scientific Contributions

Scientists at the Indian Institute of Tropical Meteorology (IITM), Pune, run advanced computer models that couple the energy flows we discussed in this chapter between atmosphere, oceans, land, and ice to simulate the Indian monsoon. These models use data from satellites, buoys in the Indian Ocean, and even stations in Antarctica to improve seasonal forecasts and to study how global warming may change monsoon rainfall patterns across India.



Fig.: Indian Institute of Tropical Meteorology (IITM), Pune

Biogeochemical Cycles

Living organisms constantly exchange matter and energy with the air, water, soil, and rocks around them. This continuous interaction between the non-living (abiotic) and living (biotic) components of the Earth results in the transfer of matter and energy across various spheres of the Earth, as discussed earlier. This process ensures that essential nutrients, such as carbon, nitrogen and oxygen, are recycled and remain available to support life on the Earth. This cyclic movement of matter and energy between the abiotic and biotic components is called the biogeochemical cycle. There is a dynamic relationship between different ecosystems. This interconnectedness helps ecosystems recover from disturbances or degradation and maintain environmental balance.

We will examine the water, carbon, nitrogen, and oxygen cycles. We will try to understand why their balance matters, and explore how they interconnect across spheres and respond to human impacts.

Water cycle

You have studied the water cycle and learnt terms like evaporation, transpiration, condensation, precipitation, infiltration and groundwater. The water cycle is summarised in Fig. In this cycle, water evaporates from water bodies, such as rivers, lakes, oceans, etc., and condenses to form clouds. It returns as precipitation in the form of rain, hail or snow to the surface and finally flows back to the ocean. Some of the water seeps through soil and rocks beneath the Earth's surface. Water dissolves minerals from soil and rocks. It also supports all terrestrial organisms and transports these nutrients to oceans, and supports marine life.



Fig.: Water cycle

Climate change is now affecting the water cycle. For example, a warmer atmosphere holds more moisture, causing heavier rains in some areas (like intensified monsoons) and droughts elsewhere. Melting glaciers add more water to rivers, raising sea levels in the long run and threatening coastal cities, such as Mumbai and Chennai. Sudden bursts of intense rainfall result in more runoff that erodes soil, and less infiltration reduces the recharge of groundwater, which in turn makes sustaining agriculture difficult, especially during dry months. In this way, the water cycle links the cryosphere (glaciers), hydrosphere (rivers and oceans), atmosphere (moisture), geosphere (soil erosion and decreased infiltration), and biosphere (crops and fisheries), all of which are affected by global warming.

Carbon cycle

Carbon forms the backbone of life. Every protein, carbohydrate, fat, and DNA molecule contains carbon. It circulates continuously between the atmosphere (CO_2 gas), biosphere (plants and animals), geosphere (carbonate rocks and fossil fuels, such as coal and oil), and the hydrosphere (dissolved CO_2 and marine shells) as shown schematically in Fig. Different parts of the carbon cycle operate at very different time scales. In the fast cycle, which happens over days to years, plants convert atmospheric CO_2 into glucose using sunlight through photosynthesis. CO_2 is released back into the atmosphere through respiration. Animals eat plants and/or other animals. When they die, CO_2 returns to the air because of decomposition.

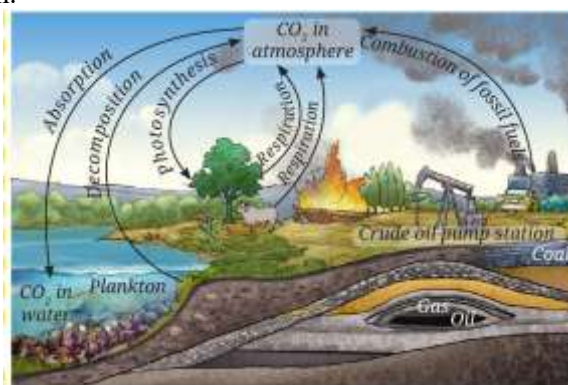


Fig.: Carbon cycle

In the slow cycle, which happens over millions of years, dead plants and animals get buried and are converted to fossil fuels, such as coal, oil and gas. These fuels are burnt to provide energy for various needs like heating, cooking, transportation and industrial purposes. When fossil fuels are burnt, carbon is released back as CO_2 on a very short time scale. The atmosphere and ocean water continuously exchange CO_2 . The ocean water absorbs atmospheric CO_2 to form carbonate and bicarbonate ions. Phytoplankton use them for photosynthesis. Some marine organisms also use them to form shells. When organisms die, they sink to the ocean floor and their organic matter is stored as carbon for a long period.

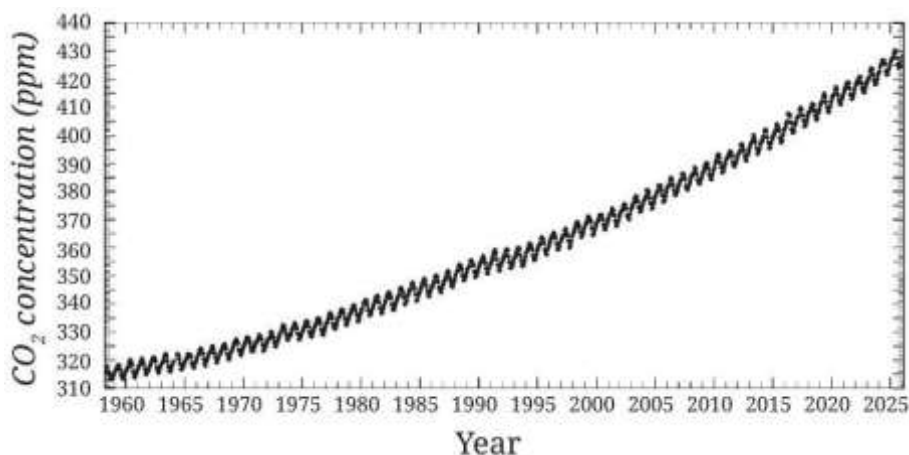


Fig.: Atmospheric CO_2 concentration (1960-2025), the Keeling curve.

The sawtooth pattern shows seasonal dips from yearly plant growth in the Northern Hemisphere absorbing CO_2

Human activities like burning fossil fuels and deforestation have raised atmospheric CO_2 by about 35% since 1960, as shown in Fig. (315 ppm to 420 ppm; ppm denotes parts per million), an unprecedented rise in the history of human civilisation. While some amounts of carbon dioxide are necessary to keep the Earth warm enough to sustain life, the balance is critical. Excessive amounts of CO_2 intensify the greenhouse effect, leading to global warming, melting of glaciers and Arctic sea ice, rising sea level, and more extreme weather conditions. In India, this may lead to more intense monsoons (warmer air holds more moisture) and threats to agriculture from its changing rainfall patterns. Though fossil fuels still power much of our electricity generation, India is rapidly increasing renewable energy sources, which will help minimise the carbon released into the atmosphere.

Threads of Curiosity

Carbon constitutes ~49% of dry weight of living organisms. Out of the total quantity of global carbon, 71% carbon is found in oceans. This oceanic reservoir regulates the amount of carbon dioxide in the atmosphere.

Do you know that the atmosphere holds only a very small fraction (about 1%) of the total global carbon?

Nitrogen cycle

Nitrogen is an essential element for the synthesis of proteins and nucleic acids in all living organisms. The largest reservoir of nitrogen is in our atmosphere. However, nitrogen gas (N_2) is rather non-reactive and cannot be directly used by plants and animals. It must first be converted to soluble compounds that living beings can absorb. The overall movement of nitrogen between air, soil, water and organisms is called the nitrogen cycle (Fig.). The nitrogen cycle contains several steps - nitrogen fixation, assimilation, ammonification, nitrification and denitrification.

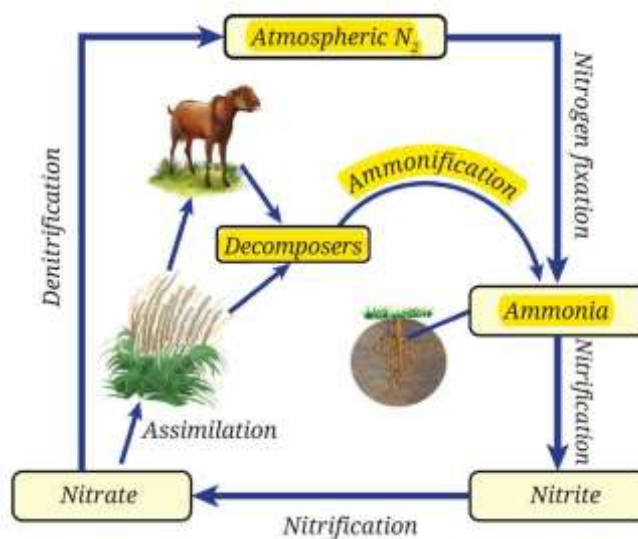


Fig.: Nitrogen cycle

Nitrogen-fixing bacteria, such as *Rhizobium* in the root nodules of legumes and *Azotobacter* in the soil, convert atmospheric N_2 into ammonia (NH_3). Nitrifying bacteria like *Nitrosomonas* convert ammonia into nitrite (NO_2^-), while *Nitrobacter* convert nitrite into nitrate (NO_3^-). This process is known as nitrification. Plants assimilate these nitrogen compounds from the soil, whereas animals obtain nitrogen by consuming plants or other animals. When plants and animals die or produce waste, decomposers like bacteria and fungi break down the organic matter, returning nitrogen compounds like ammonia to the soil. This process is known as ammonification. Denitrifying bacteria, such as *Pseudomonas*, convert some nitrates back into nitrogen gas. This process is known as denitrification. This completes the cycle and maintains a balance of nitrogen in ecosystems. Lightning also contributes to the fixation of nitrogen oxides.

Ready to Go Beyond

During lightning strikes, a tiny amount of atmospheric nitrogen is fixed to produce nitrogen oxides. Today most of the nitrogen is fixed artificially via the Haber-Bosch process (early 1900s) of making ammonia from atmospheric nitrogen, which produces most of the fertilisers used today. This 'Bread from Air' revolutionised agriculture, enabling India's Green Revolution and feeding billions. More than half the nitrogen atoms in the human body come from the Haber-Bosch process. However, this reaction is energy intensive (uses ~1–2% of global energy), and the overuse of fertilisers has degraded soil and water.

Oxygen cycle

Oxygen is one of the Earth's most abundant elements. About 21% of the atmosphere consists of free oxygen gas (O_2). It is an essential component of most biological molecules like carbohydrates, proteins, nucleic acids and fats. Oxygen also exists in combined forms - in the Earth's crust as metal oxides and minerals, and in the air as carbon dioxide. While discussing the oxygen cycle, we focus on processes that regulate the oxygen level in the atmosphere (Fig). Organisms like plants and animals use oxygen for respiration, and release CO_2 . Combustion of fuels uses oxygen and releases CO_2 . Plants restore oxygen through photosynthesis using sunlight, water and CO_2 to form glucose and release O_2 . This balance between consumption (respiration and combustion) and production (photosynthesis) circulates oxygen between the atmosphere, land, oceans and living organisms, sustaining life across all spheres of the Earth.

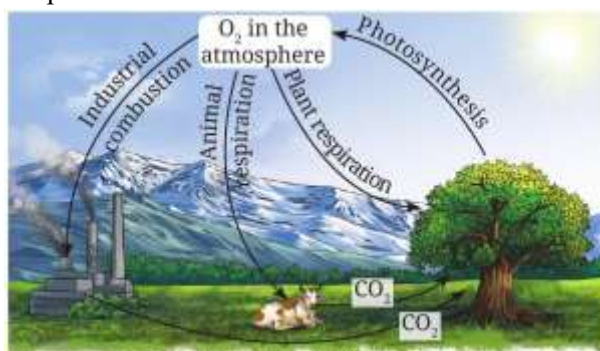


Fig.: Oxygen cycle

Human Impact on Earth's Processes

You have learnt about the impact of climate change earlier, in which we examined the crises our planet Earth is facing. Rising CO_2 levels from the use of fossil fuels lead to extreme weather and biodiversity loss. Here, we will examine how human actions disrupt the biogeochemical cycles across the Earth's spheres. Excess atmospheric CO_2 increases ocean absorption, making seawater more acidic. This could threaten tiny plankton and coral reefs, disrupting marine ecosystems. However, warmer ocean water reduces the ocean's capacity to absorb CO_2 as an effective carbon sink.

Burning of fossil fuels and deforestation saturate natural carbon sinks like forests and oceans. Over the past 50 years, there has been a significant change in the Earth's natural systems. In India, fossil fuels still generate a significant part of our electricity, contributing harmful emissions, although the rapid growth of solar energy offers hope. The excess CO_2 intensifies greenhouse warming, disrupting the carbon cycle.

The overuse of fertilisers in agriculture adds excessive nitrogen via nitrates to rivers and lakes, causing widespread growth of algae (algal blooms) that deplete oxygen and kill fish. This process is called eutrophication (Fig.), which threatens water bodies and coastal fisheries. Deforestation also creates multiple effects across air, water, land, and living organisms. Clearing forests results in decreased photosynthesis and reduced transpiration, which can lead to a decline in local rainfall. It also alters surface albedo. Without tree roots to hold the soil together, soil erosion could increase. Over time, habitats could be destroyed, leading to a decline in biodiversity as many species lose their natural homes.



Fig. Eutrophication (algal bloom)

Threads of Curiosity

Mission LiFE (Lifestyle for Environment), an India-led global initiative introduced at the United Nations Climate Change Conference in 2021, encourages people to adopt mindful, eco-friendly lifestyles. Traditional practices and ancient texts in India have long recognised that the Earth functions as an interconnected system. The Earth works through the flow of energy and the cycling of matter across spheres. Unsustainable consumption disturbs this balance. By promoting simple habits, such as saving energy and conserving resources, Mission LiFE highlights how the actions of individuals and communities can help build a sustainable future.

Vehicular emissions react with sunlight to form ground-level smog. This also leads to the formation of ground-level ozone, which is harmful to health (remember that ozone in the stratosphere blocks UV radiation and is protective of life). These pollutants make city air unhealthy.

While human impact affects all Earth's spheres, a combination of local actions and global cooperation can help restore the Earth's natural systems. For example, the Montreal Protocol has started the process of recovery of the ozone layer through global cooperation. However, the Kyoto Protocol and the Paris Agreement, in which countries were supposed to reduce their CO₂ emissions, have been less successful.

Conserving energy and energy resources, switching to renewable energy resources (like solar and wind), planting trees, saving water and practising sustainable farming can help restore the balance. India has planted billions of trees, expanded energy from solar and renewable sources significantly, and also promoted sustainable farming practices. Individuals can also contribute by saving resources (water, food and energy). This can be done by reducing waste, reusing and recycling materials. Together, all these efforts will help maintain the environmental balance on Earth.